



The impact of major warming at 14.7 ka on environmental changes and activity of Final Palaeolithic hunters at a local scale (Orawa-Nowy Targ Basin, Western Carpathians, Poland)

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Abstract

There is a widespread belief that the abrupt warming at 14.7 ka had a profound impact on the environment. However, the direct correlation between the global climatic event and changes in local environments is not obvious. We examined faunal succession in an intra-mountain basin of the Western Carpathians to assess the potential influence of the climatic change between Greenland Stadial-2a and Greenland Interstadial-1e on the local environment. We investigated three vertebrate assemblages (total number of identified specimens = 18,745; minimum number of individuals = 7515; 138 taxa) from Obłazowa Cave (western entrance) and a Rock overhang in Cisowa Rock, radiocarbon dated to the period before and after the global warming, between ca. 17.0 and 14.0 ka. Our data revealed that the major abrupt warming that occurred 14.7 ka had little impact on the local environment, which could suggest that ecosystems in Central Europe were resilient to the abrupt global climate changes. The increase in fauna population sizes and species diversities in local biotopes was gradual and began long before the temperature increase. This was supported by the analysis of ancient DNA of *Microtus arvalis*, which showed a gradual increase in effective population size after 19.0 ka. The results of palaeoclimatic reconstruction pointed out that the compared sites were characterized by similar climatic conditions. According to our calculations, the differences in the annual mean temperatures did not exceed 0.5 °C and mean

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annual thermal amplitude changed from 22.9 to 22.4 °C. The environmental changes before 14.7 ka had no impact on the activity of Final Palaeolithic hunters in the studied area.

Keywords Palaeoenvironmental reconstruction · Faunal succession · Obłazowa Cave · Magdalenian · Global warming · Climate change

Introduction

Environmental conditions are frequently considered a key component of the mode of life of Palaeolithic people, including both Neanderthals and anatomically modern humans. As a result, understanding changes in environments occupied by humans has become a main objective of prehistoric research. This knowledge is essential for reconstructing the subsistence strategies and abilities that allowed hunter-gatherers to adapt to changing environments. Pleistocene climate oscillations are thought to have greatly affected plant and animal communities re-organizing biotopes and whole ecosystems (Jackson and Blois 2015; Maguire et al. 2015); these environmental changes, in turn, also impacted past human societies.

Over the last years, studies have shown that the climate structure and changes that occurred between cold Greenland Stadials (GS) and warm Greenland Interstadials (GI), as documented in Greenland ice cores, particularly those from the last glacial period, had great millennial variabilities and reproducible patterns at the global scale (Wolff et al. 2010). In the Northern Hemisphere, in every cycle, the onset of warming events was very rapid and occurred in only a few decades, and mean annual temperature changes were even as high as between 8 and 15 °C (Huber et al. 2006).

The whole period of successive warming, sometimes called the post-LGM (post-Last Glacial Maximum), lasted about 6.3 ka, from around 18.0 ka to the beginning of the Holocene at 11.7 ka, and was characterized by two major rapid warming phases during the deglaciation period, and has been documented in both marine and terrestrial ecosystems (e.g. Allen et al. 1999; Fletcher et al. 2010; Moreno et al. 2014; Rasmussen et al. 2014). The shifts between the cold and mild periods were, as in previous climatic pulses, also very fast, as exemplified by the climatic change ca. 14.7 ka between the end of the Pleniglacial (Greenland Stadial-2a, i.e. GS-2a, an equivalent of the Oldest Dryas) and the beginning of the Late Glacial (Greenland Interstadial-1e, GI-1e, an equivalent of Bølling interstadial), when temperatures increased rapidly by 3–5 °C across western Europe (Renssen and Isarin 2001). The duration of the GS-2/GI-1 climatic change in Greenland lasted only less than a decade (Steffensen et al. 2008). This climatic shift is considered to be accompanied also by global environmental changes (Blockley et al. 2012; Clark et al. 2012; Shakun et al. 2012; Feurdean et al. 2014, among others).

It is therefore interesting to assess the influence of the rapid and large increase in temperatures, as well as of the humidity regime in the Northern Atlantic region and Greenland on particular plant and animal communities, e.g. in Central Europe. However, the impact of these global climatic changes, detected in the Greenland ice cores, on local environments as well as the correlation between discontinuous archaeological events and these geographically distant, only indirectly connected climatic regions are not inevitably applicable. The potential influence of the global climate on a given environment can be altered by the specific conditions and diversities of local landscapes and biotopes. Therefore, to understand human-environment interactions and vulnerability of local ecosystems to climate changes, it is essential to carry out detailed analyses of local palaeoenvironmental conditions in the context of global climate shifts.

To check the potential impact of global climate change on a local environment around the GS-2/GI-1 transition, we investigated succession in very rich faunal communities (amphibians, reptiles, birds and mammals) documented at two sites in north Central Europe: Obłazowa Cave (western entrance) and a Rock overhang in Cisowa Rock, both of which are in the same intra-mountain basin, the Orawa-Nowy Targ Basin, Western Carpathians, Poland. The sites include valuable deposits that accumulated continuously from ca. 17.0 to 14.0 ka. Palaeontological studies of vertebrates for archaeological and palaeoenvironmental purposes usually classify fauna into two categories, based on the size of preserved organism remains: (I) the megafauna, including large mammals, important for evaluating human subsistence strategies and patterns, and (II) the so-called small vertebrates, containing amphibians, reptiles, small bird species, insectivores, bats, rodents and lagomorphs, which can be used to reconstruct the palaeoenvironment. Small mammals, particularly rodents, have specific environmental and climatic requirements, especially relating to temperature and moisture gradients, available plant types and soil types in the areas they dig their burrows. Because of their high reproductive rates, small mammals can respond quickly to environmental changes, and this combined with their restricted territories and short migration distances, makes the group a very good indicator of the environment at local and/or regional scales (e.g. Hernández Fernández 2006; López-García et al. 2010; Socha 2014; Royer et al. 2016; Berto et al. 2017, among many others). Knowledge of small vertebrate assemblages at given time points makes it possible

to reconstruct past environments even during short climatic oscillations. Moreover, an additional advantage of small mammal assemblages is their abundance and richness, which enables the application of multiple statistical approaches and quantitative methods. To understand the population dynamics of small mammals around the GS-2/GI-1 transition, we additionally carried out genetic analysis of the common vole (*Microtus arvalis*), one of the predominant species in the assemblages, to reconstruct changes in its effective population size. Analyses of ancient DNA provide an independent information on population dynamics.

Fossil avian assemblages also supply useful environmental data at local scales; however, birds have high daily and seasonal migration abilities that must be taken into account. In the case of migrant species, the only reliable proof of their long-term presence in a given area is finding evidence of nesting and reproduction. Thus, palaeoenvironmental reconstructions based on bird fossils are particularly useful when remains of either sedentary species or young individuals of migratory birds are present in the faunal assemblages.

Amphibians and reptiles are commonly considered good indicators of past climates, although in higher latitudes there occur species adapted to the cold climate, and having wide ranges of occurrence. That is why they were not used in palaeoecological research in this work.

We also examined the link between the community structures of small mammals and birds, reflecting environmental changes at the local scale, with the activity of Final Palaeolithic hunters that inhabited the Western Carpathians.

Studied area

Remnants of Palaeolithic settlements in the Polish Western Carpathians are scarce, with the exception of the Podhale region, a southern part of the Orawa-Nowy Targ Basin (Valde-Nowak 1991). The basin has a complex tectonic history and is ca. 65 km long and up to 15–20 km wide, covering an area of ca. 300 km². It was formed during the Miocene and is partly filled with Neogene lacustrine deposits covered by large Pleistocene fluvioglacial fans of gravels, up to ca. 100 m thick (Tokarski et al. 2012). The basin is an intra-montane depression, surrounded from the west, north and east by the Flysch Carpathians and from south by the Tatra Mountains (Łanczont et al. 2019).

In the Podhale region, three neighbouring sites, Nowa Biała, site 1, Obłazowa Cave (also known as Nowa Biała, site 2) and the Rock overhang in Cisowa Rock, have been discovered in recent decades, all located in or close to the gorge of the Białka River, a tributary of the Dunajec River (Valde-Nowak 1987; Valde-Nowak et al. 2003; Nadachowski and Valde-Nowak 2015; Valde-Nowak et al. 2018; Łanczont et al. 2019) (Fig. 1). The sites are in the lower part of

Podhale, at elevations of ca. 650–700 m a.s.l., in the Pieniny Klippen Belt zone, with characteristic rocky hills that emerge in some places over a flat landscape. There are river terrace deposits in the neighbourhood of Obłazowa and Cisowa Rocks originating from the last glacial (Vistulian, Weichselian) (Baumgart-Kotarba 1996). The most important sites in this study became the multi-layered Palaeolithic Obłazowa Cave and Rock overhang in Cisowa Rock. Both sites are rich in vertebrate remains and are only 1900 m away from each other (Fig. 2). Nowa Biała 1 site did not contain animal remains (Łanczont et al. 2019).

Excavation sites

The Rock overhang in Cisowa Rock (RCR)

The Rock overhang in Cisowa Rock (49° 26' 10" N, 20° 06' 13" E) is near Gronków, at the base of the rock (685 m a.s.l.) and is completely filled with sediments (Fig. 3a). The cavity was discovered in 2009 and a regular excavation campaign was undertaken in 2013. A trench (5 × 5 × 0.6 m) was dug near the rocky wall and the whole sediment was wet sieved to discover any small teeth and bones. Preliminary studies of the faunal assemblages from layers III and I/II, yielded abundant fossil fauna with a predominance of rodents characteristic of steppe-tundra environments of the last glacial (Nadachowski and Valde-Nowak 2015); their age was confirmed by radiocarbon dating provided in this paper. The fossil remains used for this study included the whole material collected in 2013.

The Obłazowa Cave, western entrance [OB(WE)]

The Obłazowa Cave is located near Nowa Biała (49° 25' 48" N, 20° 09' 36" E), on the orographic left bank of the Białka River (Fig. 3b). The cave is in the southern part of Obłazowa Rock (688 m a.s.l.), which belongs to the Pieniny Klippen Belt, which consists of many separate rocks (klippe) that form characteristic elements of the landscape. The site is first of all known from the spectacular discovery of a boomerang made from a mammoth tusk (Valde-Nowak et al. 1987; Valde-Nowak 1991) as well as the oldest anatomically modern human (*Homo sapiens*) remains found in Poland (Gleń-Haduch 2003; Trinkaus et al. 2014). Obłazowa Cave yielded a number of important Mousterian, Szeletian, Aurignatian, Micoquian, Pavlovian and Magdalenian archaeological assemblages (Valde-Nowak et al. 2003; Valde-Nowak and Nadachowski 2014; Valde-Nowak et al. 2018) as well as a diversified fauna of molluscs, amphibians, reptiles, birds and both small and large mammals of the Late Pleistocene age, all discovered in a series of sediments divided into 21 layers (Valde-Nowak et al. 2003).

Fig. 1 View from Oblazowa Rock to the gorge and valley of the Białka River, in Podhale region, Western Carpathians, Poland. Photo. Piotr Jamiński



In 2009, a small trench was set up in the front of the western entrance to Oblazowa Cave [OB(WE), Fig. 3c]. Nadachowski and Valde-Nowak (2015) described part of the animal remains recovered during the field exploration in 2009. Direct dates obtained from bones of wild horse and reindeer place the faunal assemblage between the younger part of the GS-2 (Oldest Dryas) and the beginning of GS-1e (Bølling). The exploration of this part of Oblazowa Cave was continued over the next years, especially in 2015 and 2016, leading to the discovery of the Late Magdalenian artefacts and a fireplace in layer II, dated to GI-1d/GI-1c₃, the boundary between Bølling and Allerød (Valde-Nowak et al. 2018). The excavation yielded numerous remains of small mammals and birds together with less abundant bones of amphibians, reptiles and large mammals. Stratigraphic correlation of sediments from layers III and II in OB(WE) to the respective layers in the main chamber of Oblazowa Cave (Valde-Nowak et al. 2003) is not clear and requires further study.

Material and methods

Fossil assemblages, taxonomic diversity and taphonomy

The whole fossil assemblage excavated in 2013 from the RCR site was studied and comprised in total 3692 identified bones and teeth, which correspond to a minimum of 2006 individuals, representing at least 53 taxa, including amphibians, reptiles, birds and various species of mammals with a distinct predominance of small-sized species (Table 1, Fig. 4, Supplementary Tables S1, S2 and S3). For the OB(WE) site, the analysed material was restricted to remains collected from an area of 4 m² (N4, N3, M4 and M3) around the Magdalenian fireplace (Valde-Nowak et al. 2018), and this was enough to find potential differences in the faunal compositions of layers III and II. The faunal assemblage in OB(WE) was far more numerous and diversified than that of RCR, and comprised

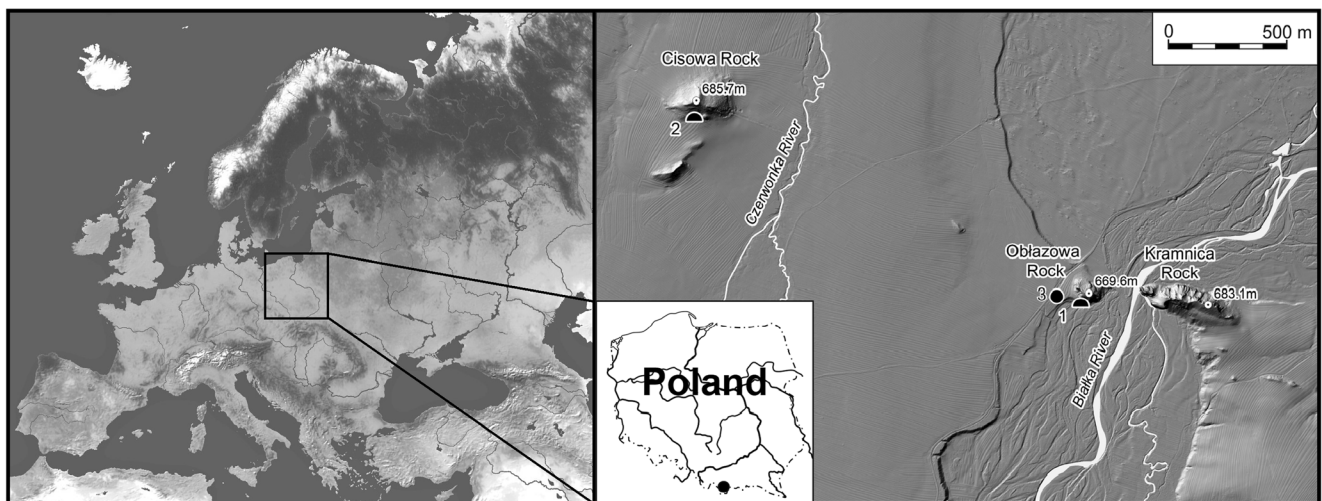


Fig. 2 Locations of Oblazowa Cave (1), Rock overhang in Cisowa Rock (2) and Nowa Biała 1 (3) sites in the Orawa-Nowy Targ Basin (Podhale region), Western Carpathians, Poland

Fig. 3 **a** Rock overhang in Cisowa Rock before excavations. **b** Oblazowa Cave, southern (main) entrance. **c** Oblazowa Cave, western entrance



15,053 identified fragments, which correspond to a minimum 5509 individuals, representing at least 120 taxa, including amphibians, reptiles, birds as well as large and small mammals (Table 2, Fig. 4, Supplementary Tables S4, S5 and S6). To estimate changes in biodiversity and compare the small mammal and bird assemblages, we used Simpson's Diversity Index (D) and Margalref's Species Richness Index (d). The statistical analyses of the distributions of individual species in the faunal assemblages were performed in R package 3.5.1. The remains were identified according to commonly used criteria, specific for every group of vertebrates, while taxonomic attributions and nomenclature followed the most recent and updated classifications. For instance, the generic name for the bank vole was changed from *Myodes* to *Clethrionomys* (Kryštufek et al. 2019), and the species name of the Late Pleistocene European narrow-headed vole was changed from *Lasiopodomys gregalis* to *Lasiopodomys anglicus*, following recent genetic studies (Baca et al. 2019). Taphonomic studies (traces of digestion and their interpretation) were based on and classified according to methods developed by Andrews (1990) and

Fernández-Jalvo et al. (2016) (the methodological, nomenclature and taphonomic details are described in the [Supplementary Material](#)).

Palaeoenvironmental analysis

In order to reconstruct the environment and landscape in the vicinity of the excavation sites, we used a method that weighted habitats according to the current distributions of small mammal species in different habitat(s) (Evans et al. 1981; Andrews 1990, 2006). The same method was applied for birds. In this approach, each species gets a maximum score of 1, which is divided among the habitat types according to its habitat preferences (described by the taxonomic habitat index, THI). Habitats were divided into the following six types according to the IUCN (2019): forests, shrublands, grasslands, wetlands, rocky areas and deserts (in this case temperate deserts). The palaeoenvironmental reconstruction of the surroundings of the sites was based on habitat frequencies (expressed as percentages). The proportions of the different

Table 1 Distribution and percentage representation of small mammals (Eulipotyphla, Chiroptera, Lagomorpha and Rodentia) in Rock overhang in Cisowa Rock

Taxon	Layer	Total (NISP/MNI)							
		III		II-III		II		I-II	
		(NISP/MNI)	%MNI	(NISP/MNI)	%MNI	(NISP/MNI)	%MNI	(NISP/MNI)	%MNI
Eulipotyphla									
<i>Sorex araneus</i>	3/2		0.49	1/1		7/3		8/5	
<i>Sorex araneus/alpinus</i>					4.00		0.30		2.67
<i>Sorex araneus/runtonensis</i>								1/1	0.53
<i>Sorex alpinus</i>								2/1	0.53
<i>Sorex minutus</i>								2/1	0.53
<i>Sorex minutus</i>						1/1	0.10	2/1	0.53
Insectivora indet.	1/1		0.23					3/2	1.07
Chiroptera									
<i>Eptesicus nilssonii</i>								1/1	0.53
Lagomorpha									
<i>Ochotona pusilla</i>	4/2		0.49						
<i>Lepus cf. timidus</i>	2/1		0.23			10/1	0.10	1/1	0.53
Rodentia									
<i>Spermophilus citelloides</i>						1/1	0.10	1/1	0.53
<i>Cricetus migratorius</i>	1/1		0.23			1/1	0.10	1/1	0.53
<i>Dicrostonyx torquatus</i>	410/251		61.01	28/15		962/545	53.85	139/73	39.04
<i>Clethrionomys glareolus</i>						2/2	0.20		
<i>Arvicola amphibius</i>						2/2	0.20	2/1	0.53
<i>Chionomys nivalis</i>	1/1		0.23	1/1		23/12	1.18	3/2	1.07
<i>Lastopodomys anglicus</i>	135/80		19.40	12/6		464/272	26.87	132/74	39.57
<i>Alexandromys oeconomicus</i>	2/2		0.49			3/3	0.30	1/1	0.53
<i>Microtus agrestis</i>	2/2		0.49			13/9	0.89	1/1	0.53
<i>Microtus arvalis</i>	119/69		16.71	4/2		318/160	15.81	33/20	10.75
Total	680/411		100	46/25		1807/1012	100	333/187	100

NISP total number of identified specimens by layer, MNI/ minimum number of individuals

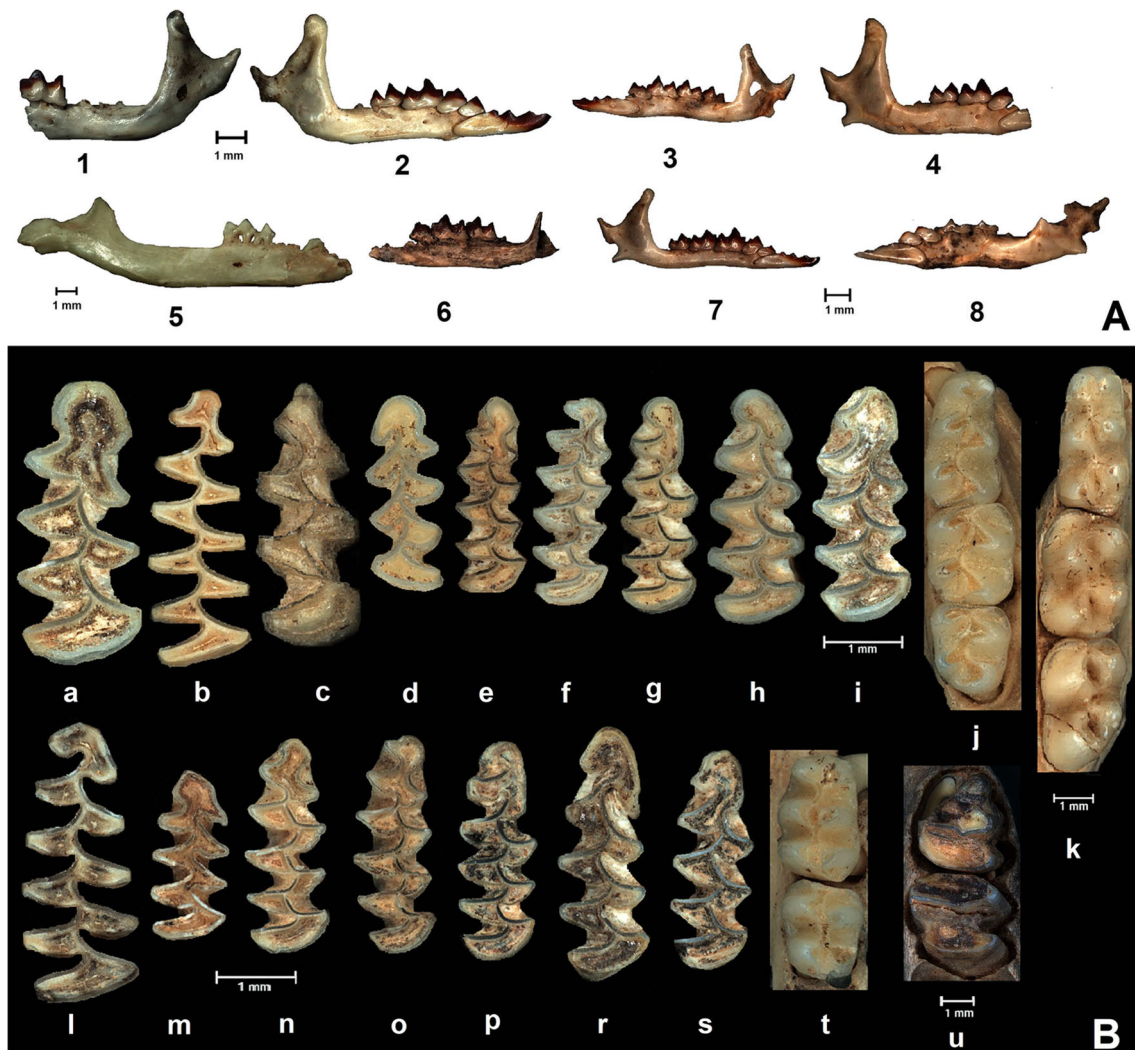


Fig. 4 Small mammal species (Eulipotyphla—A, Rodentia and Lagomorpha—B) from Oblazowa Cave, western entrance [OB(WE)] (1–5, a–k, u) and Rock overhang in Cisowa Rock (RCR) (6–8, l–t). 1—*Neomys fodiens*; 2 and 6—*Sorex araneus*; 3 and 7—*Sorex minutus*; 4—*Sorex runtonensis*; 5—*Talpa europaea*; 8—*Sorex alpinus*; a—*Arvicola*

amphibius; b and l—*Dicrostonyx torquatus*; c—*Lemmus lemmus*; d and m—*Clethrionomys glareolus*; e and n—*Microtus arvalis*; f and o—*Microtus agrestis*; g and p—*Lasiopodomys anglicus*; h and r—*Chionomys nivalis*; i and s—*Alexandromys oeconomus*; j and t—*Cricetulus migratorius*; k—*Cricetus cricetus*; u—*Ochotona pusilla*

habitats at particular sites were expressed as the sum of the given habitat preferences of individual species (Supplementary Tables S7 and S8).

Palaeoclimatic analysis

It is possible to evaluate local climatic conditions on the basis of the taxonomic composition of the small mammal assemblage assuming that there is a significant correlation between climate and mammal communities. In order to assess climate conditions in the vicinity of the RCR and OB(WE) sites, we used the qualitative and quantitative bioclimatic models for rodent faunas established by Hernández Fernández (2001a, b, 2006) and developed by Hernández Fernández and Peláez-Campomanes (2003, 2005). This method is based on

studies of 50 extant sites located in 10 climatic conditions according to Walter's (1970) typology. For each species, the Climatic Restriction Index was established: $CRI = \frac{1}{n}$, where n is the number of climatic zones inhabited by the species. If the species does not occur in a given climate type, the index is assumed to be 0, and when the species occurs in only one climate type, the index takes the value 1 (Hernández Fernández 2001a, b) (Supplementary Table S9). Based on the CRI values, we derived bioclimatic indices (BC) reflecting the proportions of the various species in each of the ten distinguished climate zones: $BC = \frac{100 \cdot \sum CRI}{S}$, where S is the number of species in the analysed assemblage. The results were used to calculate the discriminant functions of the greatest probabilities of occurrence of the assemblages per climate type (bioclimatic qualitative analysis), and the values of the BC

Table 2 Distribution and percentage representation of small mammals (Eulipotyphla, Lagomorpha and Rodentia) in Obłazowa Cave (western entrance)

Taxon	Layer				Total (NISP/MNI)
	III		II		
	(NISP/MNI)	%MNI	(NISP/MNI)	%MNI	
Eulipotyphla					
<i>Talpa europaea</i>	2/1	0.03	7/7	0.39	9/8
<i>Sorex araneus</i>	56/32	1.02	106/56	3.13	162/88
<i>Sorex araneus/runtonensis</i>	32/16	0.51	47/26	1.45	79/42
<i>Sorex minutus</i>	3/2	0.06	1/1	0.06	4/3
<i>Sorex runtonensis</i>	13/7	0.22	24/12	0.67	37/19
<i>Sorex</i> sp.	13/9	0.29	19/13	0.73	32/22
<i>Neomys fodiens</i>	4/3	0.10	2/1	0.06	6/4
Lagomorpha					
<i>Ochotona pusilla</i>	13/7	0.22	32/17	0.95	45/24
<i>Lepus</i> cf. <i>timidus</i>	257/9	0.29	159/5	0.28	416/14
Rodentia					
<i>Sciurus vulgaris</i>			2/2	0.11	2/2
<i>Spermophilus superciliosus</i>	1/1	0.03	1/1	0.06	2/2
<i>Glis glis</i>	1/1	0.03	6/4	0.22	7/5
<i>Sicista betulina</i>	3/2	0.06	4/4	0.22	7/6
<i>Cricetulus migratorius</i>	13/7	0.22	7/4	0.22	20/11
<i>Cricetus cricetus</i>	17/9	0.29	19/10	0.55	36/19
<i>Lemmus lemmus</i>	1/1	0.03	2/2	0.11	3/3
<i>Dicrostonyx torquatus</i>	1842/944	30.08	618/317	17.71	2460/1261
<i>Clethrionomys glareolus</i>	12/6	0.19	16/8	0.44	28/14
<i>Arvicola amphibius</i>	42/26	0.83	36/18	1.01	78/44
<i>Chionomys nivalis</i>	181/96	3.06	118/68	3.80	299/164
<i>Lasiopodomys anglicus</i>	1663/848	27.03	923/475	26.55	2586/1323
<i>Alexandromys oeconomus</i>	108/66	2.10	101/55	3.07	209/121
<i>Microtus agrestis</i>	182/132	4.21	113/74	4.13	295/206
<i>Microtus subterraneus</i>	1/1	0.03	3/2	0.11	4/3
<i>Microtus arvalis</i>	1744/912	29.07	1205/608	33.97	2949/1520
Total	6204/3138	100	3571/1790	100	9775/4928

NISP total number of identified specimens by layer, MNI minimum number of individuals

index for each of the 10 climatic zones were used to reconstruct 11 climatic indices using multiple linear regression; for details, see Hernández Fernández and Peláez-Campomanes (2003, 2005).

Genetic analyses

We analysed ancient DNA to estimate the demographic history of the common vole (*Microtus arvalis*), a dominant species in all our fossil assemblages. We used mtDNA sequences obtained from specimens from analysed sites to reconstruct changes in effective female population size (N_{ef}) over time. The effective population size corresponds to the number of mating individuals in an idealized population that behaves in terms of genetic drift in the same way as the actual population

analysed. This parameter is commonly used to track demographic changes in populations (Mourier et al. 2012).

Experimental work was performed in a laboratory dedicated to working with ancient DNA at the Centre of New Technologies, University of Warsaw, Poland. Sixty-three mandibles with molars or isolated molars of the common vole (*Microtus arvalis*) were used for the genetic analyses. Nineteen samples were selected from layer II of RCR; 21 and 20 specimens were selected from layers III and II of OB(WE), respectively (Supplementary Table S10). Additionally, teeth of three present-day common voles found in owl pellets collected in the vicinity of the RCR were analysed. Before DNA extraction, samples were rinsed with sterile water and crushed in sterile tubes. DNA from single rodent teeth was extracted according to Dabney et al. (2013)

and converted in double-indexed sequencing libraries using a modified protocol of Meyer and Kircher (2010). Hybridization capture was performed following Horn (2012) using mtDNA of present-day common (*Microtus arvalis*) and field voles (*Microtus agrestis*) as bait. Enriched libraries were sequenced on a NextSeq Illumina platform using a 2×75 cycles mode. After sequencing, reads were trimmed and collapsed using AdapterRemoval (Lindgreen 2012). The reads were mapped to the *M. arvalis* mtDNA sequence (NC_038176.1) and only those reads with a mapping quality over 30 and longer than 30 bp were retained. Duplicates were removed; variants and consensus sequences were called using samtools and bcftools (Li et al., 2009). The majority of samples had sequence alignments with regions where two different sequences mapped to the same reference sequence. This was most probably a result of sequencing of nuclear sequences of mitochondrial origin (pseudogenes/*numts*) alongside real mitochondrial sequences (Triant and DeWoody 2008). For further analysis, we chose a 4.3 kb long fragment between positions 12,000 and 16,285, a region to which only 1 sequence mapped in all samples. To investigate the phylogenetic position of the studied individuals and to determine to which common vole mtDNA lineage they belong, we reconstructed a Bayesian phylogeny using 303 mtDNA cytochrome *b* haplotypes (1050 bp) from present-day common voles from various regions of Europe (Supplementary Fig. S1).

To reconstruct changes in N_e of *M. arvalis*, we used the Bayesian SkyGrid method implemented in Beast 1.8.4 and Tracer 1.7 (Drummond et al. 2012, Rambaut et al. 2018). The temporal breadth of the analysed post-LGM samples was too small to reliably estimate the mutation rate in the analysed mtDNA fragment; thus, we recalculated the phylogeny using sequences from three present-day specimens and used the resulting mutation rate in the final calculation. Detailed procedures are described in the [Supplementary Material](#) section.

Radiocarbon dating

Radiocarbon dating of ten samples was carried out in Poznań Radiocarbon Laboratory (Poz) (Goslar et al. 2004) and Beta Analytic (Beta) Miami, using accelerator mass spectrometry. Taxonomically identified bone fragments or teeth were selected for the dating. In one case, a charcoal piece from the fireplace was also dated (Valde-Nowak et al. 2018). The extracted bone collagen in Poznań Radiocarbon Laboratory was ultrafiltered on pre-cleaned Vivaspin 15 MWCO 30 kD filters (Bronk Ramsey et al. 2004). There was enough collagen for dating; the quality of collagen was assessed on the basis of C/N atomic ratio and was in all cases rather low, but acceptable (interval of acceptance: 2.7–3.5; Table 3). Measurements of %C and %N were always high (39.8–51.6% and 14.7–18.9%,

respectively). The obtained radiocarbon ages were calibrated in OxCal 4.2.4 (Bronk Ramsey 2017) using the IntCal13 radiocarbon atmospheric calibration curve (Reimer et al. 2013).

Results

The vertebrate fauna of the Rock overhang in Cisowa Rock is sufficiently diverse and abundant to document faunistic changes in the studied profile. Radiocarbon dates obtained for layer II indicated that fauna accumulated during GS-2, a cold phase (ca. 17.0–16.3 ka) (Table 3; Fig. 5). The fossil assemblages preserved in layers III and II (Table 1, Supplementary Tables S1, S2 and S3) were generally similar and were combined in further analyses; however, the species composition of the uppermost layer I differed from that of the deeper layers. Layer I contained high proportions of recent humus and loose big particles of limestone rubble with small amounts of sediment filling the gaps between them. Therefore, there is a high probability that the fossil assemblage was contaminated by younger remains of the Holocene or recent age. Most probably, several species of the identified forest birds (*Coccothraustes coccothraustes*, *Chloris chloris* and *Turdus* cf. *philomelos*) (Supplementary Table S2) and some highly represented species of rodents (e.g. *Microtus arvalis* and *Clethrionomys glareolus*) (Table 1) did not belong to the Late Pleistocene assemblage. Because of the uncertain stratigraphic situation and the observed inconsistency in the species composition, the fauna from layer I was excluded from further analyses. The fauna found in layers III and II was dominated by tundra-steppe and boreal species, a vertebrate community characteristic of the end of the Late Pleistocene. During the deposition of layers III and II, the landscape surrounding the site showed periglacial features. It was mainly open, partly covered by wetlands and probably had scarce patches of shrubs or small trees. In general, the RCR small mammal fauna was distinctly less diverse ($D = 0.60$ and $d = 1.79$, Supplementary Table S11) than that of OB(WE).

The faunal assemblages of layers III and II in OB(WE) were much more diverse and abundant than in those from RCR (Table 2, Supplementary Tables S4, S5 and S6). In total, the number of determined specimens in layer III (NISP = 8948, MNI = 3422) was higher than in layer II (NISP = 6105, MNI = 2087), while the number of taxa was almost the same (ca. 89 and 86 taxa, respectively). According to AMS radiocarbon dating, layer III was older and accumulated at the end of GS-2a, between ca. 16.1 ka cal BP and 15.7 ka cal BP (95.4% range), while layer II mainly during GI-1e (Bølling), between ca. 14.7 ka cal BP and 14.0 ka cal BP, although the median of the start date (ca. 14.9 ka cal BP) for layer II falls just before the 14.7 ka cal BP warming

Table 3 Radiocarbon dates from Oblazowa Cave (western entrance) [OB(WE)] and Rock overhang in Cisowa Rock [RCR]. C/N (in coll.)—atomic ratio of C and N in extracted collagen

Site	Layer	Lab code	¹⁴ C years BP	Cal. years BP			Material Dated	C/N (in coll.)	Reference
				95.4% range		Median			
OB(WE)	II	Poz-75018	12,020 ± 90	14,120	13,704	13,881	Charcoal (fireplace)	–	Valde-Nowak et al. 2018
OB(WE)	II	Beta-519869	12,160 ± 30	14,162	13,928	14,051	<i>Cricetus cricetus</i> mandible	3.3	This paper
OB(WE)	II	Beta-517740	12,330 ± 40	14,628	14,099	14,288	<i>Rangifer tarandus</i> fragm. of naviculocuboid bone	3.2	This paper
OB(WE)	II	Poz-38123	12,340 ± 80	14,815	14,055	14,376	<i>Equus ferus</i> navicular bone	2.7	Nadachowski and Valde-Nowak 2015
OB(WE)	II	Beta-517742	12,600 ± 40	15,155	14,727	14,972	<i>Microtus arvalis</i> mandibles	3.3	This paper
OB(WE)	III	Beta-517741	13,200 ± 40	16,044	15,702	15,865	<i>Lepus</i> cf. <i>timidus</i> pelvis	3.2	This paper
OB(WE)	III	Poz-32453	13,210 ± 80	16,140	15,614	15,877	<i>Rangifer tarandus</i> medial phalanx	2.8	Nadachowski and Valde-Nowak 2015
RCR	II	Poz-88367	13,660 ± 70	16,770	16,229	16,471	<i>Lepus</i> cf. <i>timidus</i> pelvis	2.8	This paper
RCR	II	Poz-98269	13,680 ± 70	16,800	16,250	16,501	cf. <i>Rangifer tarandus</i> molar	2.7	This paper
RCR	II	Poz-98268	13,780 ± 80	16,965	16,360	16,663	<i>Equus ferus</i> bone fragment	2.7	This paper

(Table 3; Fig. 5). In general, the small mammal assemblage in layer III was less diverse ($D = 0.75$, $d = 2.86$) than that in the younger layer II ($D = 0.50$; $d = 3.20$), although species of layer III were more equally distributed than in layer II (Supplementary Table S11).

Taphonomic analysis

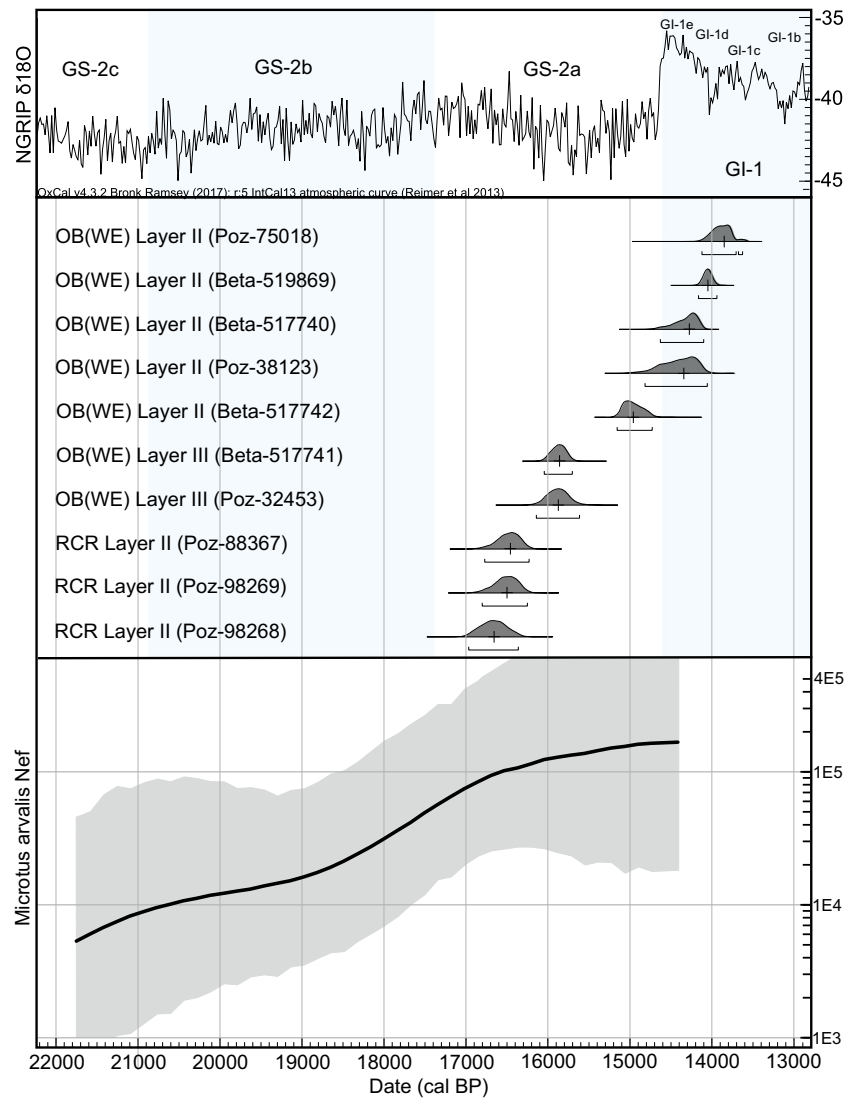
In the analysed material from the Rock overhang in Cisowa Rock, traces of digestion were found on 43% of incisors and 49% of first lower molars (Supplementary Table S12). Most of the traces belonged to the light and moderate categories. The percentage of small mammal teeth from RCR showing digestion marks and their degree of alteration point towards a category 3 predator as defined by Andrews (1990). These results indicate two possible nocturnal birds of prey, the eagle owl (*Bubo bubo*) or tawny owl (*Strix aluco*). Snowy owl (*Bubo scandiacus*), which according to Royer et al. (2019) also belongs to the third category of predators, could also be a potential predator. Remains of *Strix aluco* were found in the sediments of this site. *Strix aluco* could have been a predator and a prey of an eagle owl. Based on the remains found, this issue cannot be clearly resolved. In contrast to RCR, there were fewer traces of etching on incisors and m1s in the material from OB (WE). In layer III, traces of digestion were found on 30.5% incisors and 10.5% m1s, while in layer II, they were found on 37% incisors and 6.7% m1s. Just like in RCR, most of the traces of digestion in OB(WE) belonged to the light category. Based on traces of digestion of incisors and m1 s, it can be assumed that the potential predator was the great grey owl (*Strix nebulosa*), which fall into the second category by Andrews (1990). The results also indicate that the accumulation of digested teeth was associated with a predator that

shows a specific prey consumption pattern. Despite this, the palaeoecological interpretation based on the relative abundance of the small mammal species is a reliable indication of the habitat where predators consumed their prey.

Comparison of faunal changes in the study area

To verify how the global warming affected local fauna, we compared changes in the percentage occurrences of small mammal species in assemblages from three successive time periods represented by RCR layers III-II, OB(WE) layer III and OB(WE) layer II. The first transition occurred from the RCR layers to OB(WE) layer III, much earlier than the global warming peak between GS-2a and GI-1e, at 14.7 ka. The second shift was from OB(WE) layers III to II, which were chronologically separated by this climatic event. Figure 6 shows the results for the species that had statistically significant differences in at least one comparison between the assemblages (Table 4). Regardless of the percentage measure, %NISP or %MNI, there were clear changes in the species distributions visible not only between layers III and II of OB(WE) but also before this time, i.e. between the periods recorded in RCR layers III-II, and OB(WE) layer III. Furthermore, the changes showed the same uniform trend over time. With the exception of the collared lemming (*Dicrostonyx torquatus*), the contribution of almost all species increased from the oldest to the youngest layers. It should be emphasized that the change is generally larger between the assemblages of RCR layers III-II and OB(WE) layer III than between layers III and II of OB(WE). The frequencies of some species, e.g. *Arvicola amphibius*, *Chionomys nivalis*, *Lasiopodomys anglicus*, *Lepus* cf. *timidus* and *Microtus agrestis*, did not show significant changes at the time of the

Fig. 5 Radiocarbon dates from the studied sites plotted against the NGRIP $\delta^{18}\text{O}$ record (upper panel). GS Greenland Stadials, GI Greenland Interstadials. Bayesian reconstruction of changes in effective population size (N_e) of the common vole (*Microtus arvalis*) (lower panel)



14.7 ka climate change but did show significant increases in the previous period.

We also expressed the changes in the faunal assemblages between the studied periods by summing the absolute differences between the percentage occurrences of individual species in the assemblages. For the small mammals preserved in RCR layers III-II and OB(WE) layer III, this measure based on %NISP and %MNI was 25.6% and 26.0%, respectively, while for the assemblages recorded in layers III and II of OB(WE) it was 13.4% and 13.0%, respectively. This implies that the faunal turnover in the period before the global warming in GS-2a/GI-1e was twice as high as during the abrupt global climate change.

We applied the same procedure to avian assemblages. Due to poor representation of individual bird species, statistically significant differences were found only in a few cases (Supplementary Tables S2 and S5). Interestingly, they are all related to differences between the avifauna found in RCR

layers III-II and OB(WE) layer III. In terms of %NISP, remains of *Gallinago media*, *Strix aluco*, *Alauda arvensis*, *Pluvialis apricaria*, *Alaudidae* indet. and *Charadriiformes* indet. were more abundant in the deposits of RCR than in those of OB(WE) layer III. In the latter site, only Galliformes indet. (of *Lagopus* size) relatively dominated. No statistically significant differences were observed between OB(WE) II and III layers in this respect. These results are consistent with the sums of the absolute differences in the frequencies of individual species between these assemblages. These measures calculated for the avifauna excavated from RCR layers III-II and OB(WE) layer III are substantially larger ($A_{NISP} = 60.2\%$, $A_{MNI} = 73.4\%$) than those calculated for the faunal assemblages found in layers III and II of OB(WE) ($A_{NISP} = 10.0\%$, $A_{MNI} = 21.0\%$).

We also observed a statistically significant change in the percentage NISP between RCR layers III-II and OB(WE) layer III for *Rana temporaria*, *Vipera berus* and *Anguis fragilis*.

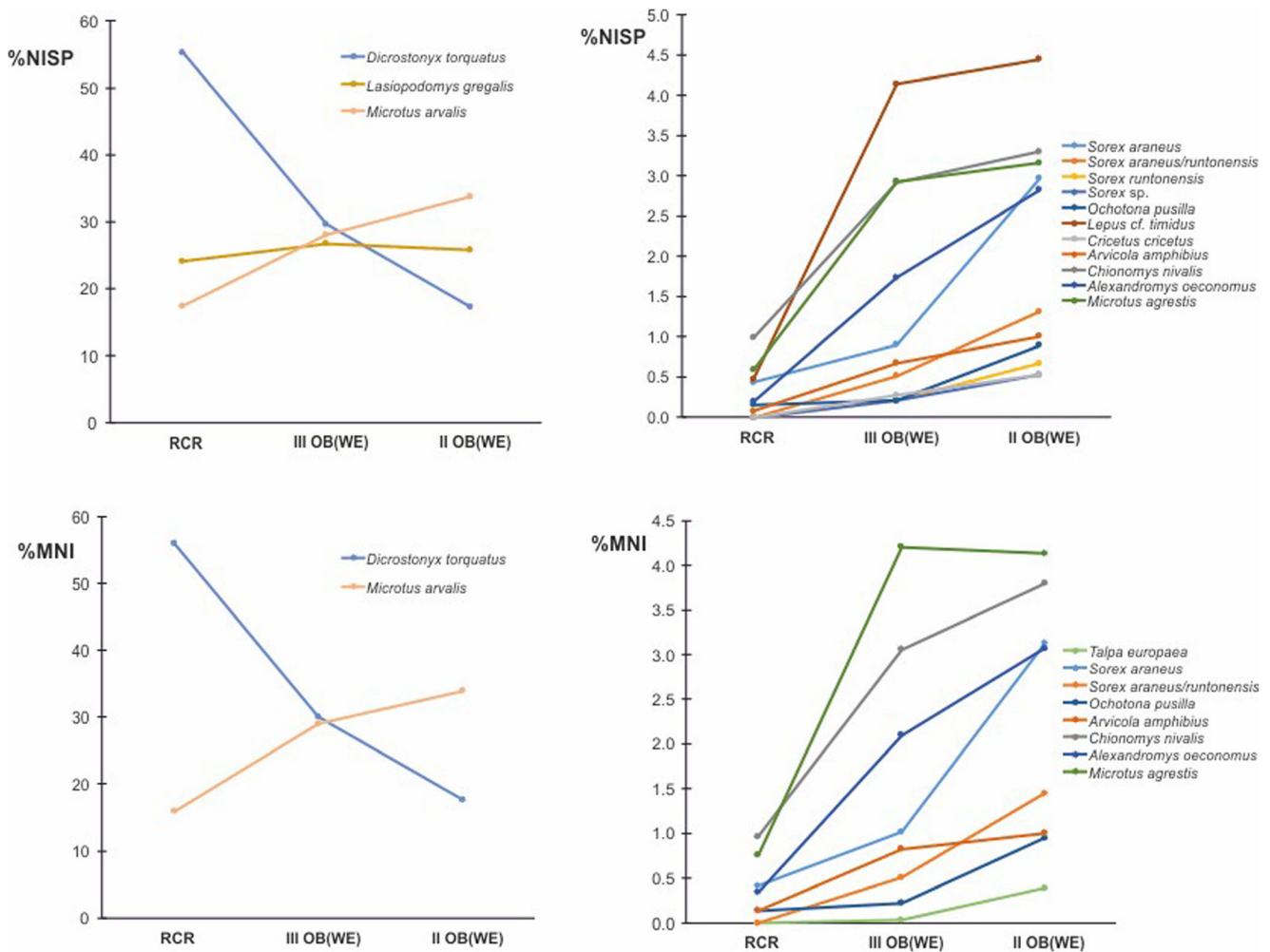


Fig. 6 Changes in the percentage occurrences (expressed in %NISP and %MNI) of individual species of small mammals in the Rock overhang in Cisowa Rock (RCR) and Obłazowa Cave, western entrance, OB(WE), layer III and layer II

The contribution of *Rana temporaria* increased with time, whereas the reptiles decreased (Supplementary Tables S1 and S4). The layers of OB(WE) did not differ significantly in terms of the distributions of amphibians and reptiles. In RCR layers III-II and OB(WE) layer III, the measures ($A_{NISP} = 54.1\%$, $A_{MNI} = 25.5\%$) were still bigger for this fauna type than those in layers III and II of OB(WE) ($A_{NISP} = 2.9\%$, $A_{MNI} = 16.1\%$).

Palaeoenvironmental reconstruction

The analyses of the landscape preferences of small mammals and birds indicated that the reconstructed environments were open areas with small patchy forests (Fig. 7a, b). There were some differences in the relative proportions of habitats between the compared sites reconstructed on the basis of small mammal and bird assemblages (Fig. 7a, b; Supplementary Tables S7 and S8). These differences seem to have been pronounced in the case of wetlands. The remains of small

mammals indicated that the share of this habitat decreased slightly over time between layers III (15%) and II (14%) of OB(WE). However, the opposite trend was observed on the basis of birds' remains. The contribution of wetlands increased from 27% in RCR layers to 29% in OB(WE) layer II. Reconstruction of the environment based on the remains of birds and small mammals showed that there was an increase in forests areas over time. The growth of forests was gradual and started before the abrupt warming at ca. 14.7 ka (Fig. 7a, b). Based on the environmental preferences of small mammals, a gradual increase in the share of forests at the expense of grasslands could be observed, although the former never outgrew the latter and equalized in the OB(WE) II layer. The share of forests increased from 25% in RCR layers to 29% in the OB(WE) II layer. At the same time, the grassland contribution dropped from 33 to 29% in the respective sites (Fig. 7a). The shrubland contribution increased from 20% in RCR layers to 22% in OB(WE) layer III and then decreased to 21% in OB(WE) layer II. Contribution of other landscapes did not

Table 4 Adjusted *p* values of the proportion test comparing the contribution of small mammal species expressed by %NISP and %MNI in three successive time periods represented by Rock overhang in Cisowa Rock (RCR) layers II to III, Oblazowa Cave, western entrance [OB(WE)] layer III and layer II. The values smaller than 0.05 were bolded

Species	RCR vs. III OB(WE)		OB(WE) III vs. II	
	%NISP	%MNI	%NISP	%MNI
<i>Talpa europaea</i>	1	1	0.065	0.034
<i>Sorex araneus</i>	0.073	0.169	3.1E-13	1.8E-06
<i>Sorex araneus/runtonensis</i>	0.002	0.051	1.7E-04	0.005
<i>Sorex minutus</i>	1	1	1	1
<i>Sorex runtonensis</i>	0.088	0.342	0.002	0.100
<i>Sorex</i> sp.	0.088	0.232	0.038	0.125
<i>Neomys fodiens</i>	0.730	1	1	1
<i>Ochotona pusilla</i>	1	1	1.9E-05	0.005
<i>Lepus</i> cf. <i>timidus</i>	3.3E-18	1	0.654	1
<i>Sciurus vulgaris</i>	NA	NA	0.460	0.437
<i>Spermophilus citelloides</i>	0.893	1	NA	NA
<i>Spermophilus superciliosus</i>	1	1	1	1
<i>Glis glis</i>	1	1	0.058	0.293
<i>Sicista betulina</i>	0.893	1	0.638	0.437
<i>Cricetulus migratorius</i>	0.523	1	1	1
<i>Cricetus cricetus</i>	0.045	0.232	0.133	0.412
<i>Lemmus lemmus</i>	1	1	0.748	0.879
<i>Dicrostonyx torquatus</i>	3.4E-110	1.26E-61	1.69E-40	3.6E-20
<i>Clethrionomys glareolus</i>	0.597	1	0.087	0.397
<i>Arvicola amphibius</i>	0.002	0.040	0.189	0.879
<i>Chionomys nivalis</i>	4.4E-07	1.3E-04	0.460	0.397
<i>Lasiopodomys anglicus</i>	0.029	0.245	0.460	0.968
<i>Alexandromys oeconomus</i>	6.4E-08	8.4E-05	0.002	0.125
<i>Microtus agrestis</i>	2.4E-10	6.4E-09	0.702	1
<i>Microtus subterraneus</i>	1	1	0.460	0.879
<i>Microtus arvalis</i>	2.0E-24	2.5E-20	4.9E-08	0.003

change. The analysis based on bird remains showed an increase in the forest contribution from 14% in RCR to 19% in OB(WE) layer II. The share of open areas increased between RCR (33%) and OB(WE) layer III (36%) and then fell in OB(WE) layer II (31%) (Fig. 7b). The contribution of shrubland decreased from 22% in RCR layers to 14% in OB(WE) layer III and then increased slightly to 15% in OB(WE) layer II. The RCR layers differed from OB(WE) layers in their greater proportions of shrubland over forest, while OB(WE) layer II was characterized by a smaller difference in the proportions of grasslands and wetlands than other layers (Fig. 7b).

The bioclimatic indices (BC) calculated for the three environments, associated with RCR layers III-II, OB(WE) layer III and OB(WE) layer II, showed the largest values for climatic zone VI (Table 5; Supplementary Table S13), characterized by a temperate climate and nemoral broadleaf deciduous forests.

Although this indicates that all three environments were similar, the detailed comparison of BC indices showed that the environments recorded in layers III and II in OB(WE) were more similar than those in RCR layers II to III and OB(WE) layer III. This indicates that the abrupt warming at ca. 14.7 ka did not influence the local environment in the studied area substantially more than in the previous period.

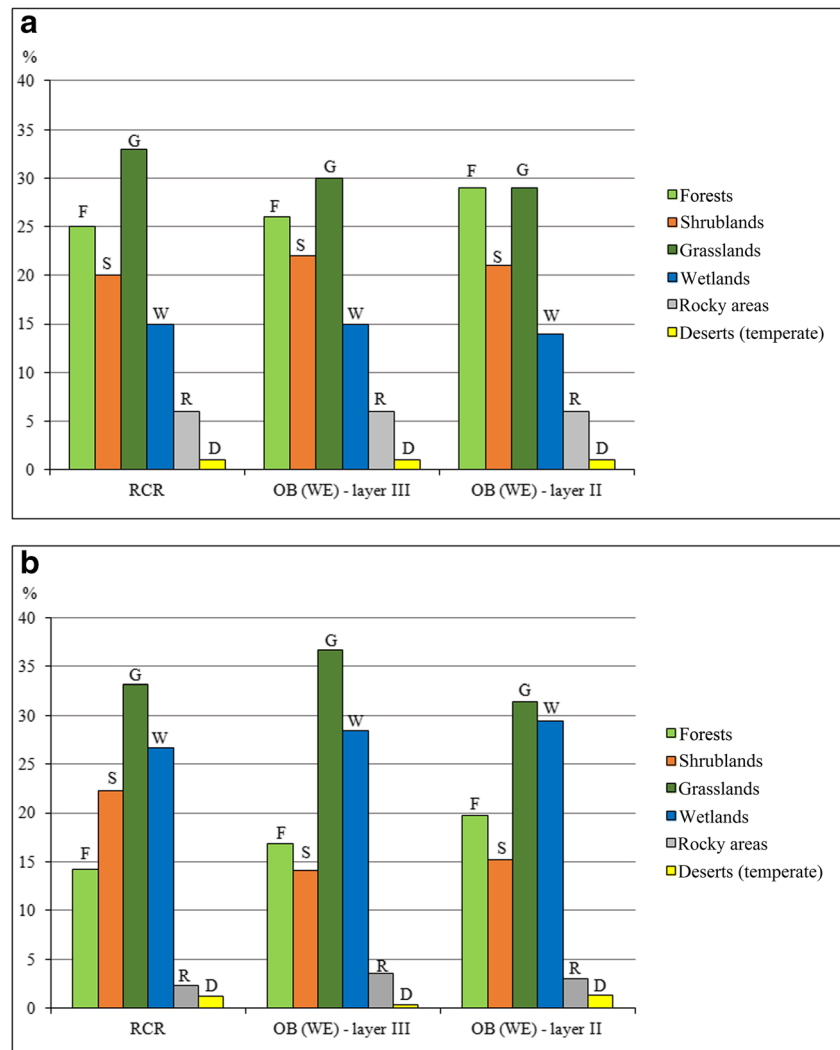
Palaeoclimatic reconstructions

The results of bioclimatic analysis based on the BC values indicated that rodents found in RCR and in layers III and II from OB(WE) lived in a typical cold temperate climate of type VI (Table 5; Supplementary Table S13). The accumulation of all compared assemblages took place under similar climatic conditions in a period characterized by relatively small temperature changes; although, as expected, many of the temperature parameters were highest for layer II in OB(WE) (Table 6). The differences in the estimated annual mean temperature (T) did not exceed 0.5 °C. The differences in the mean temperatures of the warmest month (Tmax) were even smaller and did not exceed 0.3 °C. The same was observed in the mean temperatures of the coldest month (Tmin) (differences did not exceed 0.6 °C), while the mean annual thermal amplitude (Mta) changed from 22.9 to 22.4 °C. However, differences in the values of annual total precipitation (P) were greater. The RCR assemblage accumulated during a period of relatively small annual total precipitation, while the OB(WE) assemblage from layer III accumulated in the period with highest total annual precipitation. The environment associated with the RCR layers was also characterized by the longest drought periods.

According to Walter (1970), the mean annual monthly temperature in the temperate, cool and humid climate of type VI is from + 7 to + 10 °C, i.e. higher than that reconstructed, which did not exceed + 2.7 °C. In this climate type, the mean annual precipitation is 600–800 mm, which is slightly lower than the estimated values (847–1012 mm) for the studied area. The reconstructed values of the mean annual monthly temperatures were closer to the temperatures characteristic for the very cool and humid climate of type VIII. In general, the obtained values corresponded to the boreal macroclimate (Rivas-Martinez and Rivas-Sáenz, 1996-2009), with annual mean temperatures (T) below + 6 °C and annual positive temperatures (Tp) (expressed as the sum of the mean monthly temperatures of months with average temperatures higher than 0 °C) less than 800 °C.

According to the bioclimatic division of the world by Rivas-Martinez and Rivas-Sáenz (1996-2009), bioclimates, thermobioclimates (thermotypes) and ombrotypes can be distinguished within each macrobioclimate. The divisions are based on the differences in precipitation patterns, temperatures and ombrothermal indices. By comparing the results obtained

Fig. 7 Habitat proportions in the vicinity of the Rock overhang in Cisowa Rock (RCR) and Oblazowa Cave, western entrance (OB (WE)), layers III and II reconstructed on the basis of small mammal (a) and bird (b) assemblages



for RCR and OB(WE) with the data presented by Rivas-Martinez and Rivas-Sáenz (1996-2009), it is possible to determine a more precise estimate of the climatic variation at the time when the faunal assemblages were deposited. The values of Mta, Tp and T (Table 6) indicated that the rodent assemblages were deposited in the studied sites during a subcontinental bioclimate within a boreal climate. The Tp values also make it possible to specify the thermotype within the climate. Reconstructed Tp values for the studied sites ranged between 680 and 800 °C and were characteristic of a thermoboreal climate (Tb). At present, the area corresponding to this thermoclimate, within the boreal macroclimate in Europe, is located between 25° and 65° E and between 51° and 57° N (Rivas-Martinez and Rivas-Sáenz, 1996-2009).

In spite of the general similarity between the inferred palaeoenvironments, there were larger differences between RCR layers III-II and OB(WE) layer III, than between OB(WE) layers III and II. For 9 of the 13 climatic factors, the differences were bigger for the former than the latter. This applies most to thermicity index, winter length, annual

total precipitation and drought length. If the global GS-2a/GI-1e warming had significantly affected the local environments, we should expect greater differences in the parameters calculated for OB(WE) layers III and II.

Genetic studies

To reconstruct the changes of the effective population size of common vole over time, we recovered a 4.3 kb fragment of mtDNA from 51 palaeontological and 3 present-day samples, morphologically determined as belonging to the common vole (*Microtus arvalis*). Forty-eight specimens were *M. arvalis*: 15, 18 and 15 of them were from layers II and III from OB(WE) and layer II from RCR, respectively (Supplementary Table S10). To validate the obtained sequences, we checked the amount of deamination at terminal nucleotides of DNA molecules. We also checked whether the ages assigned to mtDNA sequences were congruent with their divergence from the root of the phylogenetic tree (Root-To-Tip analysis). As a result, we discarded seven sequences as probable modern

Table 5 Values and distribution of bioclimatic index (BC) for Rock-shelter in Cisowa Rock (RCR) and Oblazowa Cave (western entrance) [OB(WE)] assemblages. Climate zones and their main vegetation types: I—equatorial zone, evergreen tropical rain forest; II—tropical zone with summer rains—tropical deciduous forest; II/III—transition tropical semi-arid zone, savanna; III—sub-tropical arid zone, sub-tropical desert; IV—

winter rain and summer drought zone—sclerophyllous woody plants; V—warm temperate zone—temperate evergreen forest; VI—typical temperate zone—nemoral broadleaf deciduous forest; VII—arid temperate zone—from steppe to cold desert; VIII—cold temperate (boreal) zone, boreal coniferous forest (taiga); IX—arctic zone, tundra

Species	Climatic zones									
	I	II	II/ III	III	IV	V	VI	VII	VIII	IX
OB (WE), layer II	0	0	0	0	7.287	0	40.618	15.100	21.868	15.100
OB (WE), layer III	0	0	0	0	5.553	0	41.106	16.106	21.106	16.106
RCR layers III and II	0	0	0	0	11.660	0	34.990	12.490	21.660	19.160

contaminants and three whose divergence from the root was substantially lower than expected at the assigned age (See [Supplementary](#) text for the details). The remaining 38 mtDNA sequences were used for reconstructing the effective population sizes. The analysis revealed a nearly 30-fold increase in the N_e between 22.0 and 14.4 ka. The N_e trajectory was not uniform across the entire period. First, there was a nearly four-fold increase between 22.0 ka and around 18.5 ka, followed by an even more rapid eight-fold increase up to a value of ca. 165,000 around 14.4 ka (Fig. 5).

Discussion

The rock overhang in Cisowa Rock and Oblazowa Cave are extraordinary sites located very close to each other, containing layers providing an uninterrupted record of faunal changes in the period from ca. 17.0 to 14.0 ka, i.e. from the end of Pleniglacial (GS-2a) to the beginning of Late Glacial (GI-1e, Bølling). This period includes an important event at ca. 14.7 ka, i.e. a global warming. Therefore, the exploration of these sites can help verify the impact of the climate change on local environments. It could be anticipated that there would be a clear relationship between the warming and the fauna preserved in the studied deposits. Specifically, as OB(WE) layer III and OB(WE) layer II correspond to the periods before and after the global warming, we should expect great differences in the faunal assemblages and bioclimatic parameters associated with these layers. Instead, we found that a faunal transition, increases in the population sizes of selected rodent species and environmental shifts proceeded continuously from ca. 17.0 ka onwards. Furthermore, we found a bigger dissimilarity in the period prior to 14.7 ka, i.e. between fauna discovered in RCR layers III-II and OB(WE) layer III, than between OB(WE) layers III and II. The changes appear to have occurred earlier, leading to an increase in biodiversity, as the small mammal (Nadachowski et al. 2003; Rzebik-Kowalska 2003) and bird (Tomek et al. 2003) assemblages discovered in

layers IV and V of Oblazowa Cave (main chamber), probably from GS-2b-a (Lorenc 2006), were generally less diverse than those found in layers III and II of OB(WE) as well as RCR. However, the proportions of species with different habitat preferences were similar in all these layers.

The presented results suggest that local climate changes in the studied region were rather unrelated to global climate changes. Correct stratigraphic recognition and dating of the layers is very important for drawing these conclusions. To dispel any doubts, we dated various fossil remains found in the layers and obtained consistent results. Moreover, we observed no disturbances relating to their deposition. We obtained five congruent, independent datings, allowing us to exclude potential contamination of the OB(WE) layer II by more recent remains of small mammals that could have indicated that the layer was younger than it really was. It could also be argued that the faunal changes occurred over time intervals of various lengths could have resulted in a greater difference between faunal assemblages in the longer period. In fact, based on the obtained radiocarbon dates (Table 3), the time interval between the average dates for RCR layers III-II and OB(WE) layer III is 674 years (89–1351 years, when considering the 95.4% range), while for layers III and II of OB(WE) it is 1557 years (459–2436 years). However, we observed larger faunal changes in the former period than in the latter. This indicates that the faunal change before the GS-2a/GI-1e boundary not only had a larger magnitude but also a faster rate.

The correctness of the stratigraphy of the studied layers is supported by conclusions drawn from other faunal assemblages in Poland of similar age. Although there are about dozen or so faunal assemblages in archaeological sites dated to the post-LGM (GS-2a) and/or beginning of Late Glacial (GI-1) (Kowalski, 1989; Madeyska and Cyrek 2002; Stefaniak et al. 2009; Wiśniewski et al. 2017), only Wilczyce (Nadachowski et al. 2014) and Krucza Skała Rock-shelter (Bocheński and Tomek 2004; Nadachowski et al. 2009) yielded sufficient remains of small mammals

Table 6 Values for the climatic factors estimated by applying quantitative bioclimatic models (Hernández Fernández and Peláez-Campomanes 2003, 2005) to rodent assemblages from Rock overhang in Cisowa Rock (RCR) and Obłazowa Cave (western entrance) [OB(WE)]

Site	T	Tp	Tmax	Tmin	Mta	It	Itc	W	VAP	FVAP	Io	P	D
OB (WE) layer II	2.7	699	14.3	− 8.3	22.6	− 139.3	− 82.3	3.1	4.6	4.8	14	987	− 0.2
OB (WE) layer III	2.5	671.4	14.0	− 8.4	22.4	− 143.9	− 91.5	3.1	4.4	4.7	14.5	1012	− 0.3
RCR layers III and II	2.2	726.6	14.0	− 8.9	22.9	− 94.7	− 94.7	3.5	4.6	4.2	13.4	847	0.4

T annual mean temperature, *Tp* annual positive temperature, *Tmax* mean temperature of the warmest month, *Tmin* mean temperature of the coldest month, *Mta* mean annual thermal amplitude, *It* thermicity index, *Itc* compensated thermicity index, *W* winter length, *VAP* vegetative activity period, *FVAP* free vegetative activity period, *Io* ombrothermic index, *P* annual total precipitation, *D* drought length

and birds to make comparisons with the fauna found in the rock overhang in Cisowa Rock and Obłazowa Cave (western entrance). Generally, the fauna described from Wilczyce and the oldest layers of Krucza Skała Rock-shelter (layers I, II and 1, 2) was very similar to that found in layers II and III of RCR as well as III of OB(WE), but was less diverse, compared especially with the latter. The differences may have resulted from local conditions, e.g. the landscape morphology, slope exposition to insolation, irrigation, soil type and species of predator responsible for accumulation of the fossil assemblages. These sites are located 150–200 km north of the sites studied here, thus, it is also possible that the lower diversity of Krucza Skała and Wilczyce was associated with their more northern locations and thus their closeness to the ice sheet and more severe conditions in the glacial period. Nevertheless, all results indicated that southern Poland was covered by steppe-tundra with a small contribution of coniferous and cold deciduous forest during GS-2a. Wong et al. (2020) came to similar conclusions by reconstructing the palaeoenvironment in Swabian Jura (southwestern Germany) based on small mammal remains. Unfortunately, the results obtained cannot be verified by palaeobotanical studies because there is no pollen profile covering a period between 17.0–14.0 ka in southern Poland.

An interesting finding of the presented analyses is that the change in faunal composition between the very cold end of the Pleniglacial (GS-2) and much warmer Bølling (GI-1e) was not as spectacular as expected. The small differences in species compositions between the compared assemblages were too small to be explained by the global climate warming at the 14.7 ka transition. Furthermore, we observed that the faunal and environmental changes started earlier, at least ca. 17.0 ka and gradually continued to ca. 14.0 ka.

The local environments may not have responded directly to the global climate changes due to the high resilience of the ecosystems (Holling 1973); this may have been associated with high biodiversity as well as high flexibility of food webs and relationships between the species in these communities. However, our findings indicated that the changes in fauna and environments began over 2000 years before and were even larger than those during the 14.7 ka global warming.

Therefore, this suggests that a climate shift could have occurred in Central Europe (or even possibly over a greater area of Europe) before the ^{18}O isotope changes were recorded in the Greenland glacier at the GS-2a/GI-1e boundary. In fact, it is believed that the main source of moisture for the ice core sites, NGRIP, GRIP and GISP2, was the North Atlantic (Charles et al. 1994; Werner et al. 2000, 2001; Langen and Vinther 2009), which is quite distant from the excavation sites analysed here.

It is not inconceivable that the climatic changes in the region that was the source of snowfall in Greenland at the end of the Pleistocene were later than those in Central Europe. It is also possible that when the fauna and environment were changing in Central Europe, the Greenland glacier was fed with snow formed by water evaporating from regions near the North Pole and thus still depleted in $\delta^{18}\text{O}$ isotope, despite the climate already being warmer in other regions. It is also possible that the successive faunal changes were not driven by temperature oscillations, but changes in other climatic parameters, e.g. precipitation and humidity, which are not directly reflected in $\delta^{18}\text{O}$ records.

The obtained result can also be explained by smaller magnitude of vegetation response to temperature increase in Central Europe (Feurdean et al. 2014) compared to the magnitude of temperature increase suggested by the Greenland oxygen isotope record at 14.7 ka (Renssen and Isarin 2001). Moreover, the magnitude of the climate shift appears to have been less dramatic in more continental Central Europe than in the more oceanic Western Europe (Feurdean et al. 2014; Moreno et al. 2014). In the period between GS-2a / GI-1e, no sudden and drastic changes in the vegetation cover were found on any of the profiles shown by Feurdean et al. (2014). Forests occurred in Central Europe before global warming 14.7 ka that confirms the obtained results. Furthermore, the Western Carpathians may have been a northern LGM refugium, especially for several species of temperate trees (Jamrichová et al. 2017), which probably resulted in the early development of suitable biotopes and increases in population sizes of several mammal species (Sommer and Nadachowski 2006) far in advance of global temperature increases. Forests have their own microclimate and a “buffer-function” that

softer climate conditions, which could be reflected by smaller and gradual changes in fauna composition pattern. Consequently, the intermediate vertebrate communities comprising a mix of extinct and extant species already existed before the period of global climate warming, whereas cold-climate species were progressively and successively replaced by temperate ones.

Final Palaeolithic in the Western Carpathians

The remains of Palaeolithic settlements in the Polish Carpathians are scarce, and the scale of research is hardly comparable to investigations conducted in other regions of Poland. In the initial phases of development of archaeology, caves in the Carpathians, even in the Tatra Mts., were treated as potential archaeological sites, e.g. Magura Cave (Jura 1955) and were studied in the same manner as those from regions nowadays considered classical for Polish archaeology, e.g. those from Kraków-Częstochowa Upland. The negative result of the Magura Cave archaeological excavation led archaeologists to abandon further research in this area. Only since the late 1970s, as a result of intensive surveys conducted in the Podhale Region and Pieniny Mts. (Valde-Nowak 1991), has the number of Palaeolithic sites, including those of Magdalenian age, increased. They represent various levels of significance. In the less important category are the surface collections in Koniówka, site 1; Ostrowsko, site 4; Frydman, site 13; Krauszów, site 1; or Szlembark, site 1 (Dryja and Rydlewski 1991; Valde-Nowak et al. 2016). Much more important are Oblazowa Cave (also known as Nowa Biała, site 2) (Valde-Nowak et al. 2003), Sromowce Wyżne-Kąty, site 1 (Valde-Nowak 1991), which are both interpreted as base-camps, and Sromowce Wyżne, a site recognized as a radiolarite workshop (Drobniiewicz et al. 1997). A separate group is most probably represented by short-term hunting camps: Dział, site 2 (Rydlewski 1990) and Podczerwone, site 1 (Valde-Nowak 1991). For Palaeolithic hunters, the Orawa-Nowy Targ Basin was a place with extremely attractive features for settlements (Łanczont et al. 2019). The basin is characterized by a varied topography and diverse geology compared with neighbouring areas, and has typical highland and even para-lowland landscapes, with easily accessible beds of good quality raw materials, especially radiolarite, which occurs in various different colour varieties. The great interest of the Palaeolithic humans in Pieniny radiolarites can be understood as a proxy indicator for the higher value of such local raw materials than, e.g. Kraków-Częstochowa flints.

The first stratified Final Palaeolithic site is Oblazowa Cave (Valde-Nowak et al. 2003, 2018). In the cave chamber, layer III, a small assemblage was found during the first excavation campaigns; in the main publication on the topic (Valde-Nowak et al. 2003), the assemblage was discussed as material probably originating from the late phase of the Upper

Palaeolithic or even Late Palaeolithic. Two blade-like end-scrapers were found. Two truncations with very fine and slightly oblique cuttings were also found and resembled Magdalenian objects (Valde-Nowak et al. 2003, 67, comp. also 57–59 Fig. 37: 4–8). As the excavations continued, some inventories typical of the Magdalenian cultural environment were discovered. The most important among them were a blade core, a Zinken type perforator, a burin and a sandstone plaque interpreted as a *Lalinde-Gönnersdorf* figurine (Valde-Nowak et al. 2017). Another Magdalenian assemblage was recently discovered under an overhang at the western entrance to Oblazowa Cave, OB(WE), in the fireplace of layer II (Valde-Nowak et al. 2018). In light of new data, we can assume that Magdalenian settlers occupied this place at the time layer III inside the cave was formed. Also noteworthy in this context is that an artefact was discovered in layer III of OB(WE) (Nadachowski and Valde-Nowak 2015). Their techno-typological features distinguish them from those discovered in the evidences in layer II (fireplace) of the overhang. However, the question of a possible connection between layer III in the chamber and layer III in the rock overhang is still open. The Palaeolithic sequence of the Oblazowa Cave chamber and western entrance ends in the older part of Allerød. The continuation of the Palaeolithic process of colonization of the vicinity of Oblazowa Rock was documented at the open air Nowa Biała 1 site, located about 100 m west of the Oblazowa Cave entrance (Łanczont et al. 2019). This site provided nearly two thousand artefacts of the Federmesser culture, one of the groups of the Arch-Backed Points Technocomplex, scattered around the remains of a hut.

Post-LGM hunter-gatherers in Sub-Carpathian Europe

The area within and around the Western Carpathians is split into three zones with different palaeogeographies, which may help us to understand post-LGM settlement in this area. The first zone is the area north of the Carpathians, which is the zone that was closest to the southern border of the Vistulian (=Weichselian) glacier and successively shifted north as the Leszno, Poznań and Pommeranian glaciation phases proceeded. Therefore, this zone did not favour permanent colonization. In fact, sites dated before the end of GS-2a are not numerous in this area. Taking into account the nature of the known traces of both Magdalenian culture in Maszycka Cave (Kozłowski et al. 2012), and Epigravettien in, e.g. Kraków-Spadzista (Kozłowski and Sobczyk 1987), Zawalona Cave in Mników (Alexandrowicz et al. 1992) and Targowisko site 10–11 (Wilczyński 2014), it has been suggested that the two cultural environments coexisted in this territory, although there is no archaeological evidence that could confirm this statement (Wiśniewski et al. 2017). In the first millennia after the LGM, the zone gives the impression that it was a “land that belonged to no one”, sporadically visited by hunter groups. The second

zone encompasses the Western Carpathians. Their northern part, i.e. the Dunajec and Poprad rivers basins, named in this article as the Orawa-Nowy Targ Basin, provide almost no traces that can be considered Epigravettian, aside from the site of Ujazd near Jasło (Valde-Nowak et al. 2006). There are many traces of the Magdalenians in Stará Ľubovňa (Valde-Nowak et al. 2007), Hłomcza on the San River (Valde-Nowak and Muzyczuk 1999), Uście Gorlickie in Lower Beskidy Mts. (Valde-Nowak 2015), as well as the Orawa-Nowy Targ Basin and Pieniny Mts, which is where the Magdalenian sites are most concentrated. These sites are of different chronology, dated from before the end of GS-2a and after. This shows that Magdalenian groups dominated in the northern part of the Western Carpathian zone. The third, southernmost zone is located in the southern part of the Western Carpathians and Carpathian Basin, to the south of the Orawa-Nowy Targ Basin. Traces of Magdalenian origin are completely absent there, and instead, Epigravettian settlements are present, e.g. at Moravany-Žakovska and Lopata (Hromada and Kozłowski 1995; Kazior et al. 1998), Sagvar (Lengyel 2016), but these sites do not show strong settlement dynamics and are very dispersed over time.

In light of these results, it can be assumed that the Orawa-Nowy Targ Basin offered favourable conditions for settlement for a long time in the Late Pleistocene, which is demonstrated by the findings in Oblazowa Cave. Therefore, the lack of unambiguous traces of Epigravettian occupation in this cave is intriguing, especially when we consider the clear presence of temporally separate remains of Magdalenian settlements. It is not inconceivable that human groups were too rare and had no opportunity to come across this region before ca. 16.0 ka or that conditions for human settlement in this area were not sufficiently attractive. Intense research in the central Western Carpathians over the last decades has shown the existence of a southern border of the Magdalenian range, which runs through the valley of the upper catchment of the Poprad River. Epigravettian sites, however few, lie south of this zone. The only Epigravettian site in the north-western Carpathians, from Ujazd, near Jasło, most probably predates the expansion of Magdalenian populations in this zone.

Conclusions

We examined the impact of the global climatic change that occurred ca. 14.7 ka, i.e. during the transition period between the Pleniglacial (GS-2a) and Bølling (GI-1e), on faunal succession at the local scale in the Orawa-Nowy Targ Basin in the Polish Western Carpathians. The study included three radiocarbon dated vertebrate assemblages. Two of them, from the Rock overhang in Cisowa Rock and layer III in Oblazowa Cave (western entrance), were deposited before the global climatic change, and one, from layer II in Oblazowa Cave (western entrance), came from the post-warming period. The

compared assemblages were very rich taxonomically, comprising altogether over 18,700 determined fossil items, representing over 7500 specimens and about 138 taxa, which allowed for reliable palaeoenvironmental analyses. We reconstructed faunal succession that evidenced a gradual environmental change occurring as early as between ca. 17.0 and 14.0 ka. Our data suggest that the major abrupt warming around GS-2/GI-1 boundary had little impact on the environment in the Polish Western Carpathians. The increase in species diversity in local biotopes was gradual and began before the general increase in temperatures in the Northern Hemisphere. Interestingly, the changes in the faunal assemblages and bioclimatic parameters were even greater in the period before the global GS-2a/GI-1e warming. This view is also supported by a gradual increase in the effective population size of the common vole (*Microtus arvalis*), starting in 19.0 ka and confirming that biotopes suitable for this species developed earlier. The vertebrate fauna found in the Rock overhang in Cisowa Rock indicated that differentiated biotopes already existed in GS-2a. The results of the palaeoclimatic reconstruction showed that the compared sites had similar climatic conditions. The differences in the mean annual temperatures did not exceed 0.5 °C, and the mean annual amplitude fell from 22.9 to 22.4 °C. The gradual environmental change also manifested in a decrease in species preferring open tundra landscapes and an increase in dwellers of patchy and more diversified biotopes. The presence of differentiated environments in the Western Carpathians at the end of GS-2a and during GI-1e correlates well with the development of Final Palaeolithic settlements; however, it is important to note that there is a lack of evidence for human presence in the area just after the LGM (between ca. 18.0 and ca. 16.0 ka) despite the presence of suitable environmental conditions for settlement. The first Late Magdalenian groups appeared in the Orawa-Nowy Targ Basin just before the major warming ca. 14.7 ka, as documented in OB(WE) layer III. However, only in the Bølling (GI-1e) was Magdalenian settlement confirmed, by findings in the main chamber of Oblazowa Cave. At the beginning of the Allerød (GI-1c₃), settlements of Late Magdalenian groups were still present at OB(WE) and other sites. The inhabitation of the region continued, as Federmesser groups appeared in the second half of this interstadial (GI-1c_{1b}) in the Nowa Biała 1 open-air campsite. It is remarkable that in the Polish Western Carpathians traces of the Epigravettian have not been found.

The main, rather unexpected result of this study is that the global warming at ca. 14.7 ka had a little influence on environmental changes in the intra-mountain basin and that the diversity and mosaicism of the local environment was already relatively high before the abrupt warming, during the post-LGM deglaciation period. These findings may imply that local ecosystems were resilient to the global climate shift. We found bigger changes in the fauna and environments prior to the 14.7 ka global warming, which suggests that changes in

$\delta^{18}\text{O}$ of Greenland ice cores were delayed compared with the transitions in climate, fauna and environment in Central Europe. However, the absence of traces of human activity in this area before ca. 16.0 ka suggests that the variety of biotopes available for humans were apparently not sufficiently attractive for settlement or the groups had not yet had the opportunity to come across this region.

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Compliance with ethical standards

Conflict of interest The authors declare that they have no conflict of interest.

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